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Factors influencing the accuracy of the denitrifier method for determining the oxygen isotopic composition of nitrate*

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Abstract: The denitrifier method is widely used as a novel pretreatment method for the determination of nitrogen and oxygen isotope ratios as it can provide quantitative and high-sensitivity measurements. Nevertheless, the method is limited by relatively low measurement accuracy for δ^{18} O. In this study, we analyzed the factors influencing the accuracy of δ^{18} O determination, and then systematically investigated the effects of dissolved oxygen concentrations and nitrate sample sizes on estimates of the δ^{15} N and δ^{18} O of nitrate reference materials. The δ^{18} O contraction ratio was used to represent the relationship between the measured difference and true difference between two reference materials. We obtained the following main results: (1) a gas-liquid ratio of 3:10 (v/v) in ordinary triangular flasks and a shaking speed of 120 r/min produced an optimal range (1.9 to 2.6 mg/L) in the concentration of dissolved oxygen for accurately determining δ^{18} O, and (2) the δ^{18} O contraction ratio decreased as nitrate sample size decreased within a certain range (1.0 to 0.1 µmol). Our results suggested that δ^{18} O contraction is influenced mainly by dissolved oxygen concentrations in pure culture, and provided a model for improving the accuracy of oxygen isotope analysis.

Key words: Denitrifier method; Nitrate; δ^{15} N; δ^{18} O; Dissolved oxygen; δ^{18} O contraction https://doi.org/10.1631/jzus.B1800197 **CLC number:** Q939.99

1 Introduction

Natural abundance stable isotope signatures have been successfully used in numerous surface water and groundwater systems to trace contamination sources (Widory et al., 2005; Popescu et al., 2015). The isotopic composition of nitrate ($\delta^{15}N$ and $\delta^{18}O$) can provide valuable information to characterize the dominant sources of nitrate pollution and quantitatively estimate the contributions from dif-

A simple, rapid, low-contamination, cost-effective, and highly precise sample preparation method for isotope measurement is urgently needed.

Continuous-flow isotope ratio mass spectrometry (IRMS) requires nitrate in water to be converted to a solid (e.g. ammonium salt or AgNO₃) and then to a gas (e.g. N₂ and CO or N₂ and CO₂). Compared with other preparation methods like ion exchange chromatography (Silva et al., 2000), the diffusion method (Stark and Hart, 1996), and chemical method (Stevens and Laughlin, 1994), the denitrifier method has the advantage of simultaneously measuring nitrogen and oxygen isotopes of nitrate at concentrations as low as 0.1 μmol/L and in sample sizes of less than

ferent sources by applying stable isotope mixing models (e.g. SIAR) (Oelmann et al., 2007; Ward et al., 2010; Rock et al., 2011; Matiatos, 2016).

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10 ml (Casciotti et al., 2002). In addition, the denitrifier method can effectively reduce errors originating from atmospheric leaks, problems associated with reduction of NO₃⁻ to NH₄⁺, or contamination from other N-sources (Christensen and Tiedje, 1988).

The denitrifier method was first introduced by Sigman et al. (2001) as a novel method to measure the natural abundance of nitrate isotopes in seawater and freshwater. It was based on the isotopic analysis of nitrous oxide (N₂O) generated from nitrate by denitrifying bacteria (e.g. Pseudomonas aureofaciens and Pseudomonas chlororaphis) lacking N₂O-reductase activity. Materials and methods like autosampler modification, water sample treatment, pre-concentration to denitrify cultures, and NO₂ removal are being constantly updated to improve the performance of the denitrifier method by increasing the precision and sensitivity of isotopic analysis (Mørkved et al., 2007; Granger and Sigman, 2009; McIlvin and Casciotti, 2011; Weigand et al., 2016). Moreover, the application of the denitrifier method is continuously expanding, and it is now used in the fields of botany, soil science, atmospheric science, and the study of the global nitrogen cycle (Dahal and Hastings, 2016).

Presently, the measurement precision of the oxygen isotope (0.5% to 2.0%) is far lower than that of the nitrogen isotope (0.20% to 0.46%) in nitrate dissolved in water samples (Vicars et al., 2013). This is because the conversion of nitrate to N_2O maintains a mass balance reaction for nitrogen but not for oxygen. Isotope fractionation and the exchange of oxygen atoms with water may lead to $\delta^{18}O$ in N_2O differing from its original nitrate.

Experiments with 18 O-labeled water were conducted by Casciotti et al. (2002). These experiments indicated that oxygen atoms exchanged with water accounted for less than 10% of those in the N₂O products for *P. aureofaciens*. In addition, both oxygen isotope fractionation and oxygen atom exchange should be consistent within a given batch of analyses. Accordingly, the effects of isotopic fractionation, exchange, and a blank on δ^{18} O measurements of water samples were accurately quantified by analyzing two reference materials (RMs) of known δ^{18} O within the same batch.

Nonetheless, due to incomplete oxygen conversion and an unclear reaction process for oxygen isotope exchange, the calibration of the oxygen isotope was complicated and the measurement precision was low

(Dai et al., 2017). Consequently, the limitation of the denitrifier method was the difficulty in accurately determining the $\delta^{18}O$ of nitrate. In addition, there was a contraction in the oxygen isotopic difference between IAEA-NO-3 (International Atomic Energy Agency, Vienna, Austria, subsequently referred to as "N3") and USUG34 (the National Institute of Standards and Technology, Gaithersburg, MD, USA) (Weigand et al., 2016).

It is known that the denitrification process is influenced by the dissolved oxygen (DO) concentration, carbon source, nitrate concentration, pH, and temperature (Deng, 2010). In addition, the experimental bacterium in the denitrifier method is an aerobic denitrifying bacterium. Therefore, the effect of DO and nitrate concentration on $\delta^{18}O$ of N_2O formed by denitrification should be noted when pH (neutral), temperature (26 °C), and carbon source are constant. In a preliminary experiment, we found that the $\delta^{18}O$ of N_2O derived from nitrate varied with the headspace in culture bottles, as noted by Xu et al. (2012).

Therefore, the objectives of this study were: (1) to examine the effects of DO concentrations and nitrate sample sizes on the precision of isotopic analysis, especially for δ^{18} O; (2) to explore the cause of poor analytical precision and the contraction of δ^{18} O; and (3) to identify the optimal DO concentrations and nitrate sample sizes to avoid δ^{18} O contraction.

2 Materials and methods

2.1 Materials

P. aureofaciens (strain number: ATCC 13985; a facultative anaerobic bacterium, recently reclassified as a strain of *P. chlororaphis*) was selected as the experimental bacterium in this study. The nitrate isotopic RMs IAEA-NO-3, USGS32, and USGS34 were used to evaluate the analytical methods according to the accuracy and precision of measurement results (Hastings et al., 2003; Weigand et al., 2016). The isotopic composition of the RMs is shown in Table 1.

2.2 Method set-up

Based on the previously published method (Sigman et al., 2001; Mørkved et al., 2007; McIlvin and Casciotti, 2011), we focused on the oxygen isotopic ratios of nitrate. We divided the procedures (Fig. 1) of

Dafamana matamial	¹⁵ N isotope ratio (×10 ³)		¹⁸ O isotope	ratio (×10 ³)	¹⁷ O isotope ratio (×10 ³)		
Reference material	$\delta^{15}N$	U ^a	$\delta^{18}O$	U ^a	$\Delta^{17}O$	U ^a	
IAEA-NO-3 ^b	4.7	±0.3	25.6	±0.4	-0.2	±0.2	
USGS32 ^c	180.0	Exact	25.7	±0.4			
USGS34 ^d	-1.8	±0.2	-27.9	±0.6	-0.1	±0.2	

Table 1 Reference values for the relative nitrogen and oxygen isotope ratios of the nitrate reference materials

the denitrifier method into the following steps: (1) resuscitation of the denitrifying strains, (2) pure culture, (3) test bottle-making, (4) sample pretreatment and nitrate conversion, (5) N_2O extraction and isolation, and (6) measurement and calibration of isotopic ratios.

First, P. aureofaciens stored in a -80 °C refrigerator was revived by daubing an agar plate. Second, a 5-ml starter was prepared from single colonies to generate inoculum overnight. The starter was transferred to a culture flask (routine lab equipment, triangular flask), sealed with a rubber stopper and Parafilm® M (#996, Pechiney Plastic Packaging, Menasha, WI, USA), and incubated using amended medium with KNO₃ for 7 d. Third, the medium was concentrated 10-fold (cell slurry was dissolved using fresh amended medium without KNO3 after centrifuging) and sealed in a 20-ml headspace vial (2 to 3 ml per vial) with butyl rubber stoppers (Ht-0942, Qingdao Hiprove Medical Technologies Co., Ltd., Qingdao, China). The headspace vial was used as the test bottle after purging for 2-3 h using helium or high purity nitrogen to remove N₂O and ensure anaerobic conditions.

Next, water samples were heated in a water bath for 1 h at 80 °C (Mørkved et al., 2007) to inhibit the activity of native denitrifying bacteria. The pH was adjusted to 3.5 using ascorbate to eliminate the disturbance of nitrite (Xu et al., 2012), and to neutral using NaOH solution. After samples (less than 10 ml) containing 0.2 to 0.4 µmol nitrate were injected into the test bottles overnight for NO₃⁻ conversion, 0.1 to 0.2 ml of 10 mol/L NaOH was used to lyse the bacteria and scrub any CO₂ gas in each test bottle. Next, N₂O was extracted from the test bottles off-line and injected into a Thermo Scientific pre-concentrator (Thermo Fisher Scientific Inc., Waltham, Massachusetts, USA) to be isolated on-line. N₂O purity was analyzed using IRMS (Thermo Finnigan MAT 253).

Finally, the RMs were processed with the samples simultaneously. The nitrate concentrations of the RMs were matched to those of the samples, and the data were used to correct the nonlinearity of the mass spectrometer and blanks associated with the procedure. The natural abundances of 15 N and 18 O were calculated as δ^{15} N and δ^{18} O, respectively, expressed in per mill units (‰): natural abundance=[($R_{\text{sample}}/R_{\text{standard}}$)-1]×1000‰, where R_{sample} and R_{standard} are the ratios of heavy isotope to light isotope of sample and standard, respectively (Liu et al., 2012). The standard for 15 N was air, and for 18 O was the Vienna Standard Means Ocean Water (VSMOW).

To facilitate analysis, we introduced the contraction ratio (C) to reflect the reliability of measurement results: $C=(T-M)/T\times100\%$, where M and T are the measured and true δ difference between RMs, respectively. Accordingly, the result is more accurate when the contraction ratio is smaller.

2.3 Experiment I: effect of DO concentrations in pure culture on δ^{18} O measurement

The DO concentrations in medium can be affected by both gas-liquid ratios in the culture bottle and shaking speeds. To explore the optimal DO concentrations in ordinary triangular flasks and the relationship between DO concentrations in pure culture and the analytical precision of δ^{18} O, we first conducted an experiment under both non-sealed and sealed conditions.

Generally, the ratio of gas to liquid was about 1:5 (v/v) for the culture bottle (Xu et al., 2012). The volume of the non-sealed pure culture with medium accounted for 5/6 of the actual volume (600 ml) of the 500 ml triangular flasks. The sealed conditions provided limited volume for the culture medium (375, 428, and 500 ml) in the flasks when other conditions were consistent with previously published methods (Sigman et al., 2001; Mørkved et al., 2007). For

^a Uncertainties for δ^{15} N, δ^{18} O, or Δ^{17} O are two times the standard uncertainty of reported values from a single study (Böhlke et al., 2003).

b Data sources (Böhlke and Coplen, 1995; Böhlke et al., 2003). Data sources (Böhlke and Coplen, 1995; Böhlke et al., 2003; Qi et al., 2003).

^d Data source (Böhlke et al., 2003)

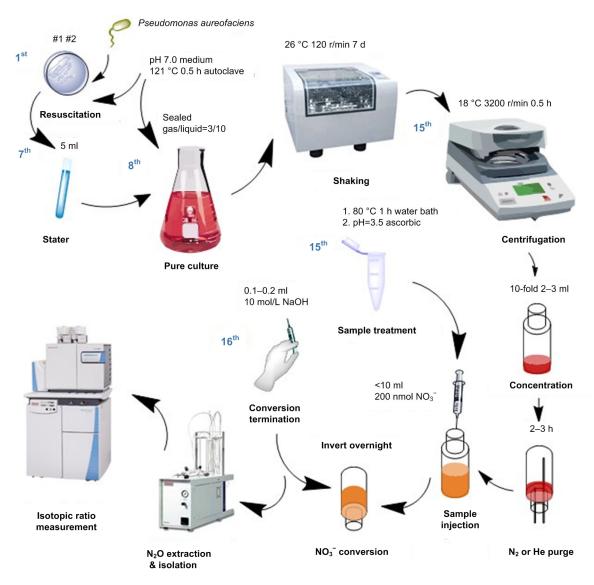


Fig. 1 Flow diagram of the experiment

non-sealed cultures, the top of each culture flask was wrapped with six layers of gauze and four layers of newspaper. For airtight flasks, rubber stoppers and Parafilm M were used to maintain predefined gasliquid ratios (3:5, 2:5, and 1:5; v/v).

After the optimal gas-liquid ratios in culture flasks were determined, two shaking speeds (60 and 120 r/min) were selected based on the literature (Sigman et al., 2001; McIlvin and Casciotti, 2011). Analytical pure (AR) KNO₃ and two RMs (USGS34 and N3) were converted to N₂O using denitrifying culture to evaluate the optimal gas-liquid ratio and shaking speed.

2.4 Experiment II: effect of nitrate sample sizes on $\delta^{18}O$ measurement

The denitrifier method has been shown to have high sensitivity (tested down to 0.01 µmol nitrate) (Casciotti et al., 2002). To examine the effect of nitrate sample size on ^{18}O measurement, the RM (N3 and USGS34) solutions were diluted across a gradient of nitrate concentrations (0.025 to 0.500 µmol/L), and then 2 ml of the diluted solution was injected into the test bottle to be denitrified in 2.5 ml of experimental bacteria solution. The analytical precision and contraction ratios of $\delta^{15}N$ and $\delta^{18}O$ in RMs were measured for different nitrate sample sizes.

3 Results

3.1 Non-sealed cultures

The average standard deviations (SDs) (n=6) of δ^{15} N were 0.25% for KNO₃ (AR), 0.28% for N3, and 0.57% for USGS34. The difference (6.38%) between measured values of δ^{15} N for USGS34 and N3 was close to the true difference of 6.5%. However, the observed variation of measured δ^{18} O values (the maximum SD=3.42%) was higher than that of δ^{15} N. The difference for measured δ^{18} O values between the two RMs ranged from 0.63% to 4.90%, which was much lower than the true difference of 53.5%. Moreover, the δ^{18} O contraction ratios were higher than 90% in non-sealed conditions.

3.2 Sealed cultures with different gas-liquid ratios

The measured δ^{18} O values of RMs under sealedculture conditions with different gas-liquid ratios are shown in Fig. 2. The results indicated that the measured δ¹⁸O differences between USGS34 and N3 increased from 7.72% to 47.98%, and δ^{18} O contraction ratios declined from 85.6% to 10.3% when the gas-liquid ratios decreased from 3:5 to 1:5 (data not shown). The contraction ratios were higher than 85.4% under the gas-liquid ratios of 3:5 and 2:5. N₂O sometimes could not be detected using IRMS when the gas-liquid ratio was 1:5 with a low cell density (optical density (OD)=0.19, while OD=0.31 for a gas-liquid ratio of 2:5). To harvest sufficient cells and guarantee complete nitrate conversion in water samples, the gas-liquid ratio in the triangular flasks needed to be greater than 1:5.

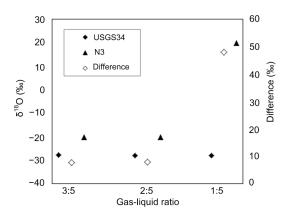


Fig. 2 Comparison of $\delta^{18}O$ values of RMs (USGS34 and N3) at different gas-liquid ratios

Therefore, the gas-liquid ratio of 3:10 (v/v) was selected for our continuous experiments, and the measurement accuracies of $\delta^{15}N$ and $\delta^{18}O$ were significantly improved. The average SD values were reduced to as low as 0.11% (N3) and 0.16% (USGS34) for $\delta^{15}N$, and 0.20% (N3) and 0.22% (USGS34) for $\delta^{18}O$. The $\delta^{18}O$ contraction ratios ranged from 4.9% to 7.3% with a mean value of 6.2%.

3.3 Nitrate sample sizes

Samples with different nitrate sample sizes were analyzed based on the case of a gas-liquid ratio of 3:10 (v/v) (Table 2). The measurement precision of δ^{18} O for either high or low NO₃⁻ sample sizes was lower than that for medium nitrate sample sizes in RMs (i.e. 0.2 and 0.4 µmol NO₃⁻, P<0.05). When the nitrate sample size was 0.05 µmol, the contraction ratios of δ^{15} N and δ^{18} O were higher than those of other nitrate sample sizes. When the nitrate sample sizes ranged from 0.2 to 0.4 µmol, the measured differences between USGS34 and N3 were close to the true difference values, and δ^{18} O contraction ratios were less than the average 6.2% (Table 2).

4 Discussion

4.1 Effect of the nitrate sample sizes on δ^{18} O measurement

The measurement accuracies of $\delta^{15}N$ and $\delta^{18}O$ were significantly improved using the gas-liquid ratio of 3:10, and the δ^{18} O contraction ratios decreased as nitrate sample sizes decreased (except for 0.05 µmol). These results agree with the observations of Weigand et al. (2016). When the nitrate sample size was 0.05 µmol, it was significantly affected by the blank, resulting in a high contraction ratio for both $\delta^{15}N$ and δ¹⁸O. Although the optimal nitrate sample sizes reduced δ^{18} O contraction ratios, there was still a slight contraction (6.0% to 6.5%) in the measured δ^{18} O values (Table 2). Weigand et al. (2016) suggested that a change in nitrate sample sizes was not the sole cause of the δ^{18} O contraction. Compared with nitrogen isotopes, oxygen isotopes are more difficult to measure due to the exchange of oxygen atoms with water, isotope fractionation, blank, and other factors.

The nitrate in water samples can be completely converted to $N_2\mathrm{O}$ in a high cell density bacterial

-	-					-	-			
No. of batches	RM	Nitrate sample size _ (μmol)	δ^{15} N (‰)		δ ¹⁸ O (‰)		Difference (‰)		C (%)	
			Mean	SD	Mean	SD	$\delta^{15}N$	$\delta^{18}O$	$\delta^{15}N$	$\delta^{18}O$
1	N3	1.00	4.59	0.08	21.57	0.27	6.41	49.62	1.4	7.3
	USGS34		-1.82	0.07	-28.05	0.33				
	N3	0.40	4.64	0.03	22.14	0.12	6.44	50.02	0.9	6.5
	USGS34		-1.80	0.19	-27.87	0.07				
	N3	0.20	4.68	0.08	22.48	0.06	6.47	50.32	0.4	6.0
	USGS34		-1.82	0.24	-27.84	0.12				
	N3	0.20	4.68	0.06	22.29	0.21	6.49	50.21	0.2	6.2
	USGS34		-1.80	0.23	-27.91	0.20				
5 N3 US	N3	0.10	4.68	0.09	23.07	0.18	6.46	50.87	0.7	4.9
	USGS34		-1.78	0.15	-27.81	0.28				
6	N3	0.05	4.54	0.29	22.32	0.36	6.20	50.12	4.7	6.3
	USGS34		-1 66	0.07	-27.80	0.30				

Table 2 Inter-batch repeatability (SD) in δ^{15} N and δ^{18} O, the measured difference between N3 and USGS34 under a gas to liquid ratio of 3:10 (v/v) in the culture flask and a shaking speed of 120 r/min (n=6), and the contraction ratio of δ^{18} O

RM: reference material; SD: standard deviation; C: contraction ratio

solution, and thus no nitrogen isotope fractionation occurs according to mass balance. However, this is not the case for oxygen atoms. Xue et al. (2010) noted that only one of the six oxygen atoms in the initial nitrate pool was represented in the N_2O (2NO₃ $^ \rightarrow$ N₂O) analysis (Fig. 3). Moreover, the intermediates (nitrite and nitric oxide; Fig. 3) during the conversion would exchange oxygen atoms with water (Ye et al., 1991). The degree of exchange may be related to the type of nitrite reductase. The copper-type nitrite reductase of P. aureofaciens incorporates fewer oxygen atoms from water into N₂O than the heme-type nitrite reductase of *P. chlororaphis* (Glockner et al., 1993). That is, not all oxygen atoms in N₂O are derived from the nitrate pool, and this issue can lead to uncertainty in isotopic analysis using the denitrifier method.

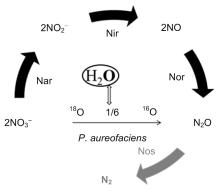


Fig. 3 Processes of denitrification by *P. aureofaciens* Nar, Nir, Nor, and Nos represent nitrate reductase, nitrite reductase, nitric oxide reductase, and nitrous oxide reductase, respectively (Morozkina and Zvyagilskaya, 2007). The gray shading in the figure indicates a process that does not occur

Moreover, there was a preferential loss of ^{16}O in these reactions that caused a difference in $\delta^{18}O$ between nitrate and the product N_2O , even if the conversion of nitrate to N_2O was complete (Casciotti et al., 2002).

The denitrifying medium may contain trace amounts of NO_2^- , NO, and N_2O generated from previous pure cultures. Previous studies showed that even if fresh medium was used instead of spent medium to re-suspend the cell slurry, and the fresh bacterial solution was purged by helium for more than 4 h, the denitrifying medium still resulted in lower blanks (about 0.04 nmol of N) (McIlvin and Casciotti, 2011). In addition, the $\delta^{18}O$ contraction of RMs may be affected by addition of KH_2PO_4 or K_2HPO_4 , the residence time of nitrite in the bacterial solution (Weigand et al., 2016), and the off-line headspace sampling (Mørkved et al., 2007).

4.2 Effect of the DO concentrations on δ^{18} O measurement

Ideally, the factors (e.g. oxygen atom exchange with water, oxygen isotope fractionation, and blank size) can be eliminated by adding RMs to a batch of samples undergoing the same operation and reaction (Casciotti et al., 2002). However, the measurement precision of δ^{18} O was low and the contraction ratios of δ^{18} O were high under non-sealed culture conditions. Different amplitudes of δ^{18} O contractions were observed under sealed-culture conditions with different gas-liquid ratios in culture flasks. That is, DO concentrations had a substantial influence on the growth of denitrifying bacteria and the conversion of nitrate

to N_2O , and this may be the most significant contributor to $\delta^{18}O$ contraction.

During denitrification, oxygen, nitrite, and nitrate are electron acceptors for nitrate conversion to N₂O in P. aureofaciens (Hu et al., 2018). If oxygen is sufficient, the bacteria can grow quickly through aerobic respiration. When oxygen drops below a certain concentration, the bacteria will switch to nitrate respiration and gain electrons from the respiratory chain (Galloway et al., 2004; Huang and Xin, 2009). McIlvin and Casciotti (2011) found that the cell density of a pure culture in an anaerobic environment was only 5% of that in an aerobic environment. It seems that the throughput of the denitrifier method can be improved using non-sealed cultures. However, higher DO concentrations may result in more disturbances to the $\delta^{18}O$ measurement of RMs so that the differences among the δ^{18} O values of RMs become indistinguishable (Fig. 2). When DO concentrations are high, nitrite reductase activity is inhibited, and the carbon source is broken down quickly to produce cells. This results in a lack of sufficient electron donors in subsequent denitrification and the accumulation of intermediate products. Therefore, appropriate DO concentrations in pure culture are essential for microbial growth and influence the completeness of nitrate conversion to N₂O.

DO concentrations in the medium depend on the dissolution rate of oxygen, gas-liquid contact time and area, gas volume, oxygen partial pressure, and the nature of the medium (Tiedje, 1988). Thus, DO concentrations were controlled by the gas-liquid ratio of culture flasks and the shaking speed under sealed conditions. Two shaking speeds (60 and 120 r/min) were tested to explore the appropriate DO concentration under a gas-liquid ratio of 3:10 (the area of gas-liquid was about 3.7 cm²). Bacterial film and flocculent precipitation were observed at a shaking speed of 60 r/min, which made it difficult to centrifuge and blend. The measured isotope ratios varied significantly, and coenobium was broken up by vortex before centrifugation in the case of 60 r/min. By contrast, the 120 r/min shaking speed allowed the growth of bacterial turbidity, but no floc, which resulted in better measurement precision of the RMs (Table 1).

Generally, denitrification is complete when the DO concentration is less than 0.2 mg/L (Downes, 1988). Most aerobic denitrifying bacteria can tolerate

DO concentrations below 3 mg/L (Zhou et al., 2007). However, nitrate may be still selected as an electron acceptor in the process of denitrification under a wider range of DO concentrations (2.3 to 11.3 mg/L) for some aerobic denitrifying bacteria (Wang et al., 2007). Our results indicated that δ^{18} O estimates of nitrate were more precise when the DO concentrations were controlled by a gas-liquid ratio of 3:10 and a shaking speed of 120 r/min. The concentrations of DO (mg/L) were 1.9 on the second, 2.6 on the fifth, and 2.2 on the seventh day of pure culture. However, the DO concentrations were more than 3 mg/L under other oxygen conditions (i.e. non-sealed culture or gas-liquid ratios of 2:5 and 3:5), which resulted in poor performance for isotopic measurement, especially for the oxygen isotope. Accordingly, the threshold values of DO concentrations for P. aureofaciens should be below 3 mg/L.

The reason why isotopic measurement was poor at DO concentrations above 3 mg/L might be that excessive DO inhibits the activity of nitrite reductase and causes the accumulation of intermediate products (Fig. 3) (Bu et al., 2015). Kumar and Lin (2010) confirmed that the reductase in the denitrification process was sensitive to oxygen, and its activity was affected by DO concentrations. A high DO concentration may promote oxygen atom exchange between intermediate products and distilled water used for the preparation of medium in the laboratory. δ^{18} O values of distilled water of less than 0% have been reported (Koehler et al., 1991). Thus, it can be inferred that the δ¹⁸O values of RMs, especially those of N3, which were lower than the true value, may be the result of oxygen atoms exchanging with distilled water. The degree of exchange was closely related to the DO concentration, and high DO concentrations may generate dilution of oxygen isotopes in intermediate products. Therefore, the effect of intermediates on the conversion of nitrate in samples to N₂O was significantly greater in high than in low DO concentrations. Eventually, the measured values of δ^{18} O in different RMs were indistinguishable due to the high DO concentration and high degree of oxygen atom exchange with water.

4.3 Precision analysis

Fig. 4 shows the results of linear regression analysis of four RMs (N3, USGS34, USGS32:USGS34,

and USGS32:2USGS34) (R^2 =1, P<0.001). Verified three times, the DO concentrations (controlled by a gas-liquid ratio of 3:10 and a shaking speed of 120 r/min) and the nitrate sample sizes (0.2 to 0.4 µmol) were effective and feasible. The slight δ^{18} O contraction can be corrected by the standard curve.

Eight rainwater samples were collected from a watershed in northwestern Zhejiang Province, China by a rainwater collector in Jan. 2017. The δ^{15} N-NO₃⁻ and δ^{18} O-NO₃⁻ values were determined with 0.2 µmol NO₃⁻. The measured results are showed in Fig. 5. The δ^{15} N-NO₃⁻ values ranged from -0.04% to 4.02% with an average SD of 0.16%, and the δ^{18} O-NO₃⁻ values ranged from 36.54% to 81.03% with an average SD of 0.25% (n=3). In general, δ^{15} N-NO₃⁻ derived from precipitation ranges from -10% to 8% (Rogers et al., 2012), and δ^{18} O-NO₃⁻ from 25% to 75% (Yang et al., 2013). Compared with the δ^{18} O precision (0.5% to 2.0%) of the water samples reported in previous studies, we obtained a slightly higher precision in this study.

5 Conclusions

In general, our study revealed a strong relationship between $\delta^{18}O$ measurements and DO concentrations or nitrate sample sizes. DO concentration in pure culture affected the subsequent conversion of nitrate to N₂O in denitrifying culture, because it promoted the accumulation of intermediate products and oxygen exchange with water. Ultimately, DO concentrations affected the measurement precision and accuracy of $\delta^{15}N$ and $\delta^{18}O$, especially $\delta^{18}O$.

In this study, we clearly demonstrated that the measurement accuracy of $\delta^{18}O$ can be improved and the $\delta^{18}O$ contraction ratio can be reduced by selecting an appropriate range of DO concentrations (1.9 to 2.6 mg/L) and nitrate sample sizes (0.2 to 0.4 µmol). The appropriate range of DO concentrations was controlled by the gas-liquid ratio of 3:10 (v/v) in ordinary triangular flasks and a shaking speed of 120 r/min. The appropriateness of the DO concentrations and nitrate sample sizes was confirmed using rainwater sample analysis. The average SDs of 0.16% for δ^{15} N-NO $_3$ and 0.25% for δ^{18} O-NO $_3$ in rainwater were slightly higher than those of water samples previously published.

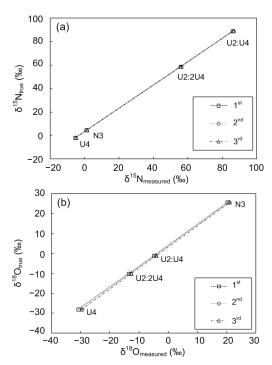


Fig. 4 Linear regression analysis of the measured values and true values for $\delta^{15}N$ (a) and $\delta^{18}O$ (b) of four reference materials

U2:U4 and U2:2U4 solutions were prepared by mixing USGS32 (U2) and USGS34 (U4) in the proportions of 1:1 and 1:2, respectively. *n*=3

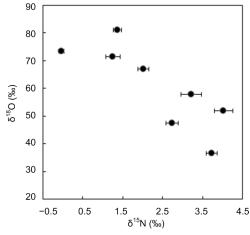


Fig. 5 Nitrogen and oxygen isotope composition of nitrate in rainwater samples (n=3)

The bars represent the standard deviations

The mechanism of the effect of DO concentrations on the exchange of oxygen with water remains to be explored, and we want to focus on the cause of δ^{18} O contraction and the determination of the concentration of oxygen isotopes.

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Compliance with ethics guidelines

Man ZHANG, Jia-chun SHI, and Lao-sheng WU declare that they have no conflict of interest.

This article does not contain any studies with human or animal subjects performed by any of the authors.

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中文概要

- 题 目:影响细菌反硝化法测量硝酸盐氧同位素的因素分析
- **目 的:** 针对硝酸盐氮氧同位素新型预处理方法——细菌 反硝化法尚存在的氧同位素测试精度低和两种 标准物质间的8¹⁸O差值会收缩等问题进行优化试 验和机理探究。
- **创新点:** 首次深入探究纯培养时的溶解氧浓度和硝酸盐进 样量对同位素测试结果的影响及原因,可为氧同 位素测试精度的提高提供方法和参考。
- 方 法: 通过控制培养瓶内的气液比(1:5、3:10、2:5和3:5) 和摇床转速(60和120 r/min)来调节纯培养时的溶解氧浓度,并设计不同的硝酸盐进样量(0.05~1.00 μmol),由硝酸盐标准物质的测试结果判断适宜的溶解氧浓度和硝酸盐进样量。
- **结** 论:标准物质间的δ¹⁸O差值收缩主要受溶解氧浓度的影响。根据氧同位素测试结果的精度和δ¹⁸O差值收缩率的大小,得出以下结论:当普通三角瓶内的气液比为3:10,摇床转速为120 r/min时的溶解氧浓度最佳(1.9~2.6 mg/L),0.2~0.4 nmol的硝酸盐进样量最适宜。
- **关键词:** 细菌反硝化法; 硝酸盐; δ^{15} N; δ^{18} O; 溶解氧; δ^{18} O 收缩